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Plant macrofossil and biomarker evidence of fen-bog transition and associated changes in vegetation

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1. Introduction

Peatlands can be divided into two main types; fens and bogs, with the main factor controlling the peatland type and the occurrence of species being the ecohydrology, i.e., the quantity and quality of water (Wheeler and Proctor, 2000; Økland et al., 2001). Fens are relatively shallow and they receive water and nutrients from the underlying and surrounding mineral soils, groundwater and atmosphere (Rydin et al., 2006). **Various** sedge species, forbs, minerotrophic Sphagna (e.g. *Sphagnum subsecundum* and *S. riparium*) as well as brown mosses such as *Warnstorfia* species **dominate fen habitats**. In contrast, due to effective peat formation and the consequential increase in **height** of the peat surface, bogs are nutrient poor as they receive water and nutrients only through precipitation, maintaining plants, including

hummock *Sphagna* (e.g. *Sphagnum fuscum*), dwarf shrubs, lichens and true mosses (e.g. *Pleurozium schreberi* and *Polytrichum* spp.) (Rydin et al., 2006). The different environmental conditions in terms of the level of acidity, nutrient status and water table level means **that raised bog peats in boreal region** usually contain less-humified peat, while in groundwater-fed less acidic fen environments biomass decay is much faster and highly humified peat layers are formed (Moore et al., 2007).

Peatlands play an important role in atmospheric carbon cycling. Northern peatlands alone are estimated to store 547 (473-621) Pg organic carbon (Yu et al., 2010), yet, simultaneously peatlands are a natural source of CH₄ to the atmosphere (Matthews, 2000). As different peatland habitats and their vegetation, even within one peatland complex, have a vital role in the carbon budget of the peatland (e.g. Riutta et al., 2007), and the fact that bryophyte- and vascular plant- dominated communities differ in their CO₂ and CH₄ dynamics (Laine et al., 2007; Levy et al., 2012), it is important to be able to separate different habitats when reconstructing peatland dynamics back in time (Yu et al., 2013). Historical peatland habitats are preserved in peat layers as decomposed plant remains that form a key proxy when reconstructing the carbon budget of the peatland or historical climatic conditions (e.g. Barber et al., 1998; Mauquoy et al., 2002; Tuittila et al., 2007; Välranta et al., 2007). Due to natural peatland succession **towards ombrotrophic conditions driven by peat height growth, minerotrophic (Frolking et al. 2010) fen peat layers are likely deposited under most of the southern boreal raised bogs.** Moreover, most of the high latitude peatlands are still the fen type of peatlands (e.g. Turunen et al., 2002) and in these environments the lack of identifiable plant remains may hamper palaeoecological and -climatological reconstructions.

Recently organic geochemistry analyses have shown that the lipid fractions of plants contain biomarkers for identifying different plant species and plant groups from peat archives.

Studies on bog peat environments have shown that plant group-specific chemical compounds can be applied to identify fossil plants or plant groups from peat (e.g. Avsejs et al., 2002; Bingham et al., 2010; Jia et al., 2008; McClymont et al., 2008; Xie et al., 2000). The most widely analyzed compounds have been the *n*-alkanes: for instance, the difference between concentrations of mid chain length (*n*-C₂₃ and *n*-C₂₅) and long chain length (*n*-C₂₉ and *n*-C₃₃) *n*-alkanes have been used to separate contributions of *Sphagnum* and vascular plant species in the peat (e.g. Andersson et al., 2011; López-Días et al., 2010; Nichols et al., 2006; Ortiz et al., 2011; Pancost et al., 2002; Ronkainen et al. 2013; Vonk and Gustafsson, 2009). Furthermore, the *n*-C₂₃/*n*-C₂₅ alkane ratio has been successfully used in tracking changes in *Sphagnum fuscum* abundance in a peat section from Finland (Bingham et al., 2010).

Thus far, in environmental reconstructions the biomarker analyses have focused on bog peats and plants typical to bogs. However, a pertinent question remaining is whether such plant-specific biomarkers could also provide information about the past plant assemblages in fen environments characterized by highly humified peat, where macrofossil remains are heavily degraded and thus essential information for environmental reconstruction is lost. Some previous studies have applied “bog” biomarker analyses throughout the peat profile, including the fen peat section underlying the bog peat section (e.g. Andersson and Myers 2012; Andersson et al., 2011). However, our recent study of the *n*-alkane concentrations, *n*-alkane ratios, and sterol distributions of moss and vascular plant species typical to fen habitats **showed** differences between some of the biomarker distributions between plant species characteristics to bogs and fens (Ronkainen et al., 2013). **As in** previous studies on bog plants (Baas et al., 2000; Ficken et al., 1998; Nichols et al., 2006; Pancost et al., 2002) fen *Sphagnum* species were also dominated by mid-chain *n*-alkanes and the above ground parts of fen vascular plants by long-chain *n*-alkanes. However, results showed similarity in the dominating *n*-alkanes of *Sphagnum* species and below-ground parts of sedges in studied

fen plants and thus, applying *n*-alkane ratios from bog plants to fen peats could result in incorrect interpretations about the actual proportions of these plant groups in peat (Ronkainen et al., 2013). Similar mid-chain *n*-alkane distributions in vascular plant roots were also reported by Huang et al. (2011). The similarity of *n*-alkane distributions in *Sphagnum* species and vascular plant below-ground parts suggest that the *n*-alkane ratios that have predicted different plant groups in bog environments relatively well (e.g. Andersson et al. 2011; Bingham et al., 2010) may not be directly applicable to interpret past habitats in environments where sedges dominate and sedge roots form an important peat component. One potential way to overcome this problem could be to combine the distributions of plant group-specific *n*-alkanes, *n*-alkane ratios and sterols, if present, when analyzing the biomarker data (Ronkainen et al., 2013). The degradation of the plant matter and their chemical compounds could impede detection of especially sterols from the fen-bog transition and fen environment, where the peat **is usually** highly humified. The level of organic matter decomposition can be studied by comparing the variations in C/N ratio and the amount of total organic carbon (TOC) through the peat section (Kuhry and Vitt 1996) In addition, the ratio between 5 α (H)-stanols and Δ^5 -sterols can be used to infer the rate of sterol degradation because 5 α (H)-stanols are known as degradation products of the Δ^5 -sterols (McClymont et al., under review). A high 5 α (H)-stanols and Δ^5 -sterols ratio is related to anoxic conditions, shallow water table level and high rate of degradation (Bertrand et al., 2012; McClymont et al., under review). The level of degradation can also be estimated by *n*-alkane CPI value, where high CPI value is linked to high amount of well-preserved plant material (Andersson and Meyers 2012; Xie et al., 2004).

In this study we analyzed fossil plant and biomarker compositions of two peat sections. We concentrated on the fen-bog transition phase where the plant composition is known to change

(e.g. Dudová et al., 2013; Loisel and Yu, 2013; Salojärvi et al. in prep.; Tuittila et al. 2013). We aimed to investigate (1) if biomarkers can be applied to distinguish fen and bog environments, and (2) if plant-specific biomarkers can be identified from fen peat. For this we applied plant macrofossil analysis to examine the past plant compositions from the same subsamples from which the selected organic geochemical-analyses were **obtained**.

2. Material and methods

2.1. Sampling

Two peat sections (SJ5 and SJ6) were collected from two closely located peatlands in Siikajoki (64°45'N, 24°42'E) located near west coast of Gulf of Bothnia, Baltic Sea in the mid-boreal bio-climate zone in Finland (Fig. 1). A chronosequence of terrestrial ecosystems from coast to inland have been created by the postglacial isostatic rising, and along the sequence peatlands alternate with sand dunes and glaciofluvial ridges (Tuittila et al., 2013). We have previously studied vegetation, microbial communities and carbon dynamics along a transect of seven mires (SJ0 to SJ6) (e.g. Leppälä et al. 2011a,b, Laine et al. 2011, Tuittila et al. 2013, Larmola et al. *in press*). In this study we concentrated on the two oldest sites of the transect, and in particular on their sediment sections where fen-bog transition occurred, from which historical plant communities had already been studied (Tuittila et al. 2013). Site SJ5 represents a peatland stage where the fen-bog transition is still partly in progress; vegetation is a mosaic of wet fen communities and drier bog communities, the average water table level is at 12 cm below moss layer surface and the basal age of the peatland is 2520 (±20) years BP. The peat core was taken from a drier surface dominated by lawn species (*S. magellanicum*). Site SJ6 is already a true bog with vegetation formed by hummock *Sphagna* and dwarf shrubs, average water table level 32 cm below moss surface and the basal age of 3000

years (both cores basal ages are extrapolated from known land-uplift rate). The top section of the peat core was dominated by hummock species (*S. fuscum*). More detailed site descriptions can be found in Leppälä et al. (2011), and Tuittila et al. (2013). The sampling depth for SJ5 was 6 – 150 cm and for SJ6 0 – 100 cm. Both cores were cut into 2 cm sample slices and analyzed with a varying down-core resolution by focusing on the fen-bog transition.

2.2. Plant macrofossil analyses

Plant macrofossil sample volume was 5 cm³. Samples were rinsed under running water using a 140-µm sieve and no chemical treatment was needed. Remains retained on a sieve were identified and the percentage in volume of constitute within the total composition of the sample was visually estimated by using a stereomicroscope (magnification of 10x) (e.g. Speranza et al., 2000; Mauquoy et al., 2002) If the proportion of bryophytes exceeded 10 % of the total sample volume a high power light microscope was used to identify bryophyte species and to count proportions for different bryophyte species. Also, the proportion of unidentified organic matter (UOM) from samples was estimated (cf. Väiliranta et al., 2007 and references therein).

2.2. Solvent extraction

Peat samples for solvent extraction were freeze dried and ground following the same procedure in Ronkainen et al. (2013). Lipids were extracted from ca. 0.2 g of samples using repeated ultrasonication (20 min) with 6 ml CH₂Cl₂/MeOH (3:1, v/v). Samples were saponified with 0.5 M methanolic (95%) NaOH for 2 h at 70 °C and the neutral lipids were extracted using hexane. The neutral lipids were further separated into apolar and polar compounds using activated Al₂O₃ columns, eluting with hexane/CH₂Cl₂ (9:1, v/v) and

CH₂Cl₂/MeOH (1:2, v/v), respectively. Prior to analysis using gas chromatography (GC) and GC-mass spectrometry (GC-MS), the polar fractions were derivatised using bis(trimethylsilyl)trifluoroacetamide (Sigma Aldrich).

2.3. GC-MS

Apolar and polar fractions were analyzed using GC-MS with a gas chromatograph equipped with flame ionisation detection (GC-FID) and split/splitless injection (280 °C). Separation was achieved with a fused silica column (30 m x 0.25 mm i.d) coated with 0.25µm 5% phenyl methyl siloxane (HP-5MS), with He as carrier gas, and the following oven temperature program: 60 – 200 °C at 20 °C/min, then to 320 °C (held 35 min) at 6°C/min. The mass spectrometer was operated in full scan mode (50-650 amu/s, electron voltage 70eV, source temperature 230 °C). Compounds were assigned using the NIST mass spectral database and comparison with published spectra (e.g. Goad and Akihisa, 1997; Killops and Frewin, 1994). Quantification was achieved through comparison of integrated peak areas in the FID chromatograms and those of internal standards of known concentration (5-α-cholestane for apolars and 2-nonadecanone for polars). Biomarker concentrations were normalized to total organic carbon (TOC) content and are presented here as concentration per g TOC, so that samples with different extent of degradation become comparable (Meyers 2003; Ortiz et al., 2010). Total organic C and N₂ were measured by the CHN elemental analysis, where 1-2 mg dried and ground sample was combusted at 950°C with He as a carrier gas. The reduction of the combustion gases was carried out in a separate furnace, and separated into individual components on a temperature programmed desorption column and fed into a thermal conductivity detector. Results were computed as concentrations of C and N₂ from the detector signal.

2.4. Statistical analysis

We applied multivariate analyses to study the variation within the plant macrofossil and biomarker data (triterpenoids, sterols, stanols, *n*-alcohols (C₂₀-C₂₈) and *n*-alkanes (C₂₀-C₃₅) (µg/g TOC), and *n*-alkane ratios, **see the supplementary data 2**). For the analyses we combined macrofossil and biomarker data from both cores by depth.

We first quantified separately the variation in macrofossil plant species and biomarkers within the peat profiles by unconstrained (indirect) gradient analysis. **For macrofossils we used detrended correspondence analysis (DCA) where detrending was conducted by segments. All identified macrofossils were included in the analysis, with down-weighting of rare species. Macrofossil data was log transformed. For biomarkers we used principal component analysis (PCA) with centered and standardized data. No exclusion of biomarkers due low concentrations was done.** Secondly, for the biomarker data we conducted redundancy analysis (RDA), a constrained (direct) gradient analysis, to test if the variation in biomarkers in the peat profiles correlates with the variation in macrofossil data. For the analysis we applied the sample scores along the first and second macrofossil DCA axes as explanatory variables. **Similarly to PCA the biomarker data was centered and standardized. All constrained axes were tested with unrestricted Monte Carlo permutation (number of permutations 499).** Multivariate analyses were conducted by using Canoco for Windows 4.52 (ter Braak and Smilauer, 2002). **The correlation of the ten most significant biomarkers identified in RDA with depth and UOM** was analysed with Pearson two-tailed correlation using SPSS PASW statistics 18.

We applied TWINSpan (Two Way Indicator SPecies ANalysis, Twinspan for Windows 2.3) to define groups of macrofossils and biomarkers that share a similar abundance peak in the peat profile . For the analysis we rescaled the abundances for each macrofossil and biomarker from 0 to 1 by setting the highest abundance of each unit to 1 and calculating other

values as a percentage of the highest abundance of the unit. In the analysis we used five cut levels (0.0, 0.2, 0.4, 0.6 and 0.8) of abundance and two division levels, which determines the maximum level of recursive splitting for samples and for species (Hill and Šmilauer, 2005). The statistical analyses allow us to assess the relationship of the biomarkers to the macrofossil record and to determinate if macrofossils and biomarkers can separate different environmental habitats as individual or rather as combined proxies.

3. Results

3.1. Macrofossil analyses

Sub-fossil plant assemblages revealed clearly the vegetation succession from fen to bog stage (Fig. 2). Core SJ5 showed dominance of higher plants (*Menyanthes trifoliata*, *Scheuchzeria palustris*, *Equisetum* sp.) in the earliest stage of the succession (150 – 100cm). In core SJ6 these plants were present but they did not dominate (100 – 80 cm). In both of the studied cores the transition from fen to bog environment is **both identified and induced** by the appearance of *Eriophorum* sp. and sedge roots followed by a distinctive occupation of *Sphagnum* mosses at the depth of 40 – 20 cm in SJ5 and 75 – 55 cm in SJ6.

The DCA for macrofossil data (cores SJ5 and SJ6 combined) sub-divided the peat samples into three different groups: fen species (SJ5 150 – 30 cm and SJ6 98 - 62 cm), lawn species (SJ5 26 – 6 cm), and hummock species (SJ6 60 – 0 cm). The first axis describing the fen-bog gradient explained **28 %** (eigenvalue **0.5298**) of the variation in the data and the second axis describing the moisture gradient on bog explained **18 %** (eigenvalue **0.3491**) of the variation (Fig. 3).

3.2. Biomarker analyses

3.2.1. Apolar fraction

Figures 4a and 4b show the distribution of *n*-alkanes in both peat cores (SJ5 and SJ6). In both cores, there is up-core variation in both overall *n*-alkane concentration and the chain-length of the dominant *n*-alkane. *n*-C₂₇ dominated the bottom layers (150 – 36 cm) of SJ5, excluding layers 130 – 110 cm which were dominated again by *n*-C₂₃ and *n*-C₂₅ alkanes, while the uppermost layers (32 – 0 cm) were mainly dominated by *n*-C₂₃ and *n*-C₂₅ alkanes. In the deepest layers (100 – 60 cm) in core SJ6, the dominant *n*-alkane alternated between *n*-C₂₃, *n*-C₂₉ and *n*-C₃₁. The middle layers (60 – 40 cm) were dominated by *n*-C₂₅ whereas the uppermost layers (30 – 0 cm) of core SJ6 were dominated by the *n*-C₃₁ alkane. Different *n*-alkane ratios showed changes along the depth in both cores (Fig. 5). In core SJ5 the ratios *n*-C₂₃/*n*-C₂₇, *n*-C₂₃/*n*-C₂₉, *n*-C₃₁/*n*-C₂₇ and *n*-C₃₁/*n*-C₂₉ showed differences along the core, separating the top and the bottom layers apart; all ratios being higher than 1 in top 25 cm. For core SJ6 several ratios were able to separate the top and bottom layers apart e.g. *n*-C₂₃/*n*-C₂₅ and *n*-C₂₅/*n*-C₂₃ indicated changes happening at 62 cm (Fig 5).

The distribution of the detected triterpenoids along the cores was similar for all compounds. In core SJ5 at depths 42 – 30 cm the maximum concentrations of taraxer-14-ene (ca. 60 – 720 µg/gTOC), and taraxast-20-ene (ca. 50 – 500 µg/gTOC) were recorded. Squalene was identified only in upper layers, peaking at 30 cm (30 µg/gTOC). In core SJ6 at depths 80 – 66 cm the maximum concentrations of taraxer-14-ene (260-1050 µg/gTOC) and taraxast-20-ene (500 – 1800 µg/gTOC) were recorded. Concentration of squalene was highest at 10 cm ca. 140 µg/gTOC, downcore the concentration was ca. 10 µg/gTOC (Fig 6). The highest concentrations of taraxer-14-ene and taraxast-20-ene coincided with high counts of sedge roots and UOM in the middle layers of both cores (Figures 2 and 6).

3.2.2. Polar fraction

247 The most abundant sterols found from both cores were: brassicasterol [(22*E*)-ergosta-5,22-
 248 dien-3 β -ol], campesterol [campest-5-en-3 β -ol], stigmasterol [(24*E*)-stigmasta-5,22-dien-3 β -
 249 ol], and β -sitosterol [(3 β)-stigmast-5-en-3-ol]. The associated stanols of these sterols were
 250 also detected: campestanol [24-methyl-5 α -cholestan-3 β -ol], 22*E*-stigmastanol [(24-ethyl-5 α -
 251 cholest-22-3 β -ol)] and 3-stigmastanol [(24-ethyl-5 α -cholestan-3 β -ol)]. In core SJ5 the
 252 concentration of brassicasterol was highest (ca. 150 μ g/gTOC) at 50 – 30 cm; below and
 253 above this depth the concentration was less than 50 μ g/gTOC, although concentrations
 254 increased in the uppermost 6 cm (80 μ g/gTOC). The concentration of campesterol was also
 255 high in the uppermost 6 cm of core SJ5, reaching concentrations of ca. 2000 μ g/gTOC.
 256 Between 50 – 30 cm depth the concentration of campesterol was ca. 500 – 1000 μ g/gTOC,
 257 and elsewhere in the core the concentration was ca. 100 – 470 μ g/gTOC. Stigmasterol in core
 258 SJ5 had similar concentrations and pattern as campesterol, and whilst β -sitosterol also
 259 followed this same pattern the concentrations were significantly higher, reaching a maximum
 260 of ca. 12 000 μ g/gTOC at 30 cm depth (Fig. 6).

261 In core SJ6 the concentration of brassicasterol was higher than in core SJ5, it increased
 262 towards the top of the core with highest concentration at top 30 cm (200 – 500 μ g/gTOC).
 263 The concentration of campesterol in core SJ6 was the highest at depths 80 – 60 and 0 cm
 264 (ca. 1250 – 1700 μ g/gTOC). Also stigmasterol peaked at 62-60 cm (1300 μ g/gTOC) and at 0
 265 cm (1900 μ g/gTOC). The concentration of β -sitosterol was higher in core SJ6, peaking at 72
 266 cm (20 000 μ g/gTOC), elsewhere concentration varied between 2500 and 10 000 μ g/gTOC.
 267 Tocopherol- α [(2*R*)-2,5,7,8-Tetramethyl-2-[(4*R*,8*R*)-(4,8,12-trimethyltridecyl)]-6-chroman-
 268 ol] was found only from the lowermost layers of core SJ6. In both cores all stanols, campestanol,
 269 22*E*-stigmastanol and 3-stigmastanol, were identified from bottom to upwards at all layers
 270 until 10 cm in SJ5 and at 40 cm in SJ6 (Fig. 6).

In core SJ5 the concentration of phytol [(3,7,11,15-tetramethylhexadec-2-en-1-ol)] was highest at 50 cm depth (ca. 2500 µg/gTOC), and the overall concentration decreased towards the top layers. In core SJ6 the concentration of phytol was highest in top layers, 60 – 0 cm, ca. 400 – 1200 µg/gTOC. Before 66 cm, phytol concentration was less than 400 µg/gTOC. The *n*-alcohol distribution in both studied cores did not vary substantially. In core SJ5 *n*-C₂₂-ol dominated depths 150, 120 and 100 – 6 cm, *n*-C₂₄-ol dominated depths 140, 130 and 110 cm, and *n*-C₂₈-ol dominated depth 70 cm. In core SJ6 *n*-C₂₂-ol dominated depths 100 – 50 cm, whereas *n*-C₂₄-ol dominated the uppermost 40 cm.

3.2.3. Statistical analyses of biomarkers

The PCA for biomarkers **produced** groups, **similar to the** DCA for macrofossils. However, when the biomarker RDA was performed (Fig.7), where sample scores from macrofossil DCA were used as explanatory variables, the biomarker distribution of the two peat profiles correlated significantly with their macrofossil compositions (pseudo $F = 9.2$, p -value = 0.002). RDA resulted in three groups similar to macrofossil DCA: fens (SJ5 150 – 30 cm and SJ6 98 – 62 cm), lawn (SJ5 26 – 6 cm), and hummock (SJ6 60 – 0 cm; Fig. 3). Biomarkers whose concentrations decreased in association with the shift from fen to bog habitat were the *n*-alkanes *n*-C₂₀, *n*-C₂₂, *n*-C₂₄, *n*-C₂₆, *n*-C₂₇, *n*-C₂₈ and stanols. Markers for the top layers of core SJ5 (lawn) were e.g. the *n*-alkane ratios *n*-C₂₃/*n*-C₂₇ and *n*-C₃₁/*n*-C₂₉, and for the top core of SJ6 (hummock) were e.g. *n*-C₂₅, *n*-C₂₉, and *n*-C₂₈-alcohol. The ten best-fitted biomarkers from **the** RDA are shown in figure 8. In core SJ5 *n*-alkanes *n*-C₂₂, *n*-C₂₄, *n*-C₂₅, *n*-C₂₆, *n*-C₂₇ correlated positively and 22E-stigmastanol negatively with depth, and only *n*-C₂₅, *n*-C₂₇, *n*-C₂₉ correlated with UOM. In core SJ6 most of the *n*-alkanes correlated positively with UOM and depth, with *n*-alkane concentrations decreasing towards top layers (Fig 8).

The first division of the TWINSpan separated macrofossils from the top part of core SJ5, i.e. lawn species with urs-12-ene from the rest of the samples (Table 1). The second division divided the rest of the data hummock species and biomarkers: (1) *S. fuscum* and shrub leaves with C₂₇ *n*-alcohol, C₃₄ *n*-alkane, and squalene, and (2) fen species and biomarkers: brown-mosses, sedge-roots, *Equisetum* sp., *Sch. palustris* and *M. trifoliata* together with biomarkers: 22E-stigmastanol, *n*-C₃₅ and urs-12-ene.

3.2.4. Degradation measures

In both cores a high amount of UOM corresponded to the fen-bog transition zone (SJ5: 30 – 20 cm and SJ6: 75 – 55 cm). In these layers 20-55 % of the plant macrofossil material was unidentified. In contrast, total organic carbon (TOC) showed little variation and stayed around 50 % throughout both of the cores (Fig 9). There was a clear up-core increase in the C/N ratio in both cores at the fen-bog transition. The most notable increase of the carbon preference index (CPI; e.g. Andersson et al. 2012) of *n*-alkanes also occurred during the transition. The ratio of 5 α (H)-stanols/ Δ^5 -sterols (Bertrand et al. 2012) decreased towards the top layers in both cores and was last detected in SJ5 at 10 cm and in SJ6 at 40 cm depth.

3. Discussion

Our results suggest that, statistically, individual biomarkers predict the fossil plant species composition rather poorly, in support of observations from previous studies (e.g. Andersson 2012; Ficken et al., 1998; Pancost et al., 2002) that recommended that biomarkers should be used as a complementary proxy. When we applied the combined data i.e. biomarkers together with macrofossils as explanatory variables a clear correlation between biomarkers and fossil plants was detected and the biomarkers succeeded in describing three different environments: bog hummocks and lawns, and fens (Fig. 7).

320 Previous studies have identified the high concentrations of $n\text{-C}_{23}$ and $n\text{-C}_{25}$ and high ratios of
 321 $n\text{-C}_{23}$ to $n\text{-C}_{25}$, $n\text{-C}_{29}$ and $n\text{-C}_{31}$, to be characteristic to hummock *Sphagnum*-species, whereas
 322 taraxer-14-ene, taraxas-20-ene, $n\text{-C}_{31}$, and the ratio of $n\text{-C}_{31}/n\text{-C}_{33}$ to *Ericaceae*-species (e.g.
 323 Bingham et al., 2010; McClymont et al., 2008; Nichols et al., 2006; Nott et al. 2000; Pancost
 324 et al., 2002). In contrast, the study here shows that the statistically significant biomarkers for
 325 bog hummock species, *S. fuscum*, *S. angustifolium* and *Ericaceae* roots and leafs were $n\text{-C}_{24}$ -
 326 ol, $n\text{-C}_{26}$ -ol, $n\text{-C}_{28}$ -ol, $n\text{-C}_{25}$ and $n\text{-C}_{29}$ (Fig 7). These compounds were particularly effective
 327 in identifying the difference between bog and fen zones (Figures 7 and 8). However, the
 328 visual comparison between the biomarker concentrations and palaeobotanical assemblages
 329 supports the previous works that **has** linked certain plant groups with certain biomarker
 330 distributions; for example in core SJ6 the uppermost layers were dominated by $n\text{-C}_{31}$, which
 331 is an indicator of *Ericaceae*-species whose macrofossils were also present. However, the
 332 triterpenoids and sterols associated with *Ericaceae*-species (e.g. Pancost, 2002) were not
 333 detected. Although *S. fuscum* dominated the whole hummock bog peat section (SJ6) the
 334 concentration of $n\text{-C}_{25}$ was **exceeded** by $n\text{-C}_{31}$ when descending from 0 cm to 40 cm, but the
 335 low ratio of $n\text{-C}_{23}/n\text{-C}_{25}$ (< 1) indicates a dry *Sphagnum*- dominated environment, as
 336 suggested by Bingham et al. (2010). In core SJ5, similar low $n\text{-C}_{23}/n\text{-C}_{25}$ ratios were detected
 337 in layers dominated by sedge roots, which agrees with Ronkainen et al. (2013), whose data
 338 showed this ratio to correspond both with sedge below-ground parts and *Sphagnum* mosses.
 339 However, in core SJ6 the depths that were dominated by sedge roots (100 – 66 cm) have a
 340 higher $n\text{-C}_{23}/n\text{-C}_{25}$ ratio than comparable layers in core SJ5 (100 – 40 cm) (Fig.5). These
 341 results support the conclusion of Ronkainen et al. (2013) who suggested that the application
 342 of bog biomarkers to fen environments may be complicated by the similar signatures of
 343 *Sphagnum* mosses and sedge roots.

344 A recent study showed that in general the most reliable proxy for *Sphagnum* mosses in peats
 345 are the *n*-alkane ratios $n\text{-C}_{23}/n\text{-C}_{27}$ or $n\text{-C}_{23}/n\text{-C}_{29}$ (Bush and McInerney, 2013). When studied
 346 visually rather than through the statistical analysis our results showed that SJ6 top peat layers
 347 (60 – 0 cm) dominated by *S. fuscum* were separated from the rest of the layers by low *n*-
 348 $\text{C}_{23}/n\text{-C}_{29}$ ratio (< 0.5). The statistically significant ratio $n\text{-C}_{23}/n\text{-C}_{27}$ (< 1.5) described core
 349 SJ5 top layers (36 – 0 cm) that were dominated by *S. magellanicum* and *S. papillosum*. Other
 350 statistically significant biomarkers describing the uppermost layers of the core SJ5 with lawn
 351 habitat were $n\text{-C}_{31}/n\text{-C}_{27}$ and $n\text{-C}_{31}/n\text{-C}_{29}$ (Fig.7 and 8). The *n*-alkanes that dominated the
 352 lawn layer with *S. magellanicum* and *S. papillosum* were $n\text{-C}_{23}$, $n\text{-C}_{25}$ and $n\text{-C}_{31}$, which agree
 353 with Bingham et al. (2010). Consistent with the fact that lawns are relatively wet
 354 microhabitats when compared **with** hummocks, the previously suggested $n\text{-C}_{23}/n\text{-C}_{25}$ and *n*-
 355 $\text{C}_{23}/n\text{-C}_{31}$ ratios that should describe dry bog environment (Bingham et al. 2010) did not
 356 describe the wetter lawn environment.

357 In both cores the fen layers beneath bog peat consisted mainly of vascular plant remains, e.g.
 358 *M. trifoliata*, *Sch. palustris*, *Equisetum* spp. and sedges. These plants are usually dominated
 359 by odd-over-even long-chain *n*-alkanes (Fig. 2) and this was at least partly shown by RDA
 360 that grouped the mid- and long-chain *n*-alkanes C_{20} , C_{21} , C_{22} , C_{24} , C_{26} , C_{27} and C_{28} as well as
 361 three stanols as fen peat biomarkers (Fig.7). Bush and McInerney (2013) stated that $n\text{-C}_{29}$ and
 362 $n\text{-C}_{31}$ should not be used as general proxies for grasses and woody plants, as these two
 363 compounds are highly variable and are overlapping between these groups, but that by
 364 differences in mid-chain and long-chain *n*-alkanes *Sphagnum* mosses could be separated from
 365 them. Our results partly agree with this. In both of the studied cores the macrofossil records
 366 indicated the transition zone from fen to bog stage (SJ5; 30 – 20 cm, SJ6; 75 – 55 cm)
 367 distinctively. In core SJ5, the biomarker record indicates that the fen-bog transition is

characterized by a shift from long-chain *n*-alkanes (C₂₇) to mid-chain *n*-alkanes (C₂₃) at depth 36 cm. Yet, in core SJ6 such a clear change is not visible.

In our study of modern fen species, we found that sterols such as lupeol [5 α -lup-20(29)-en-3 β -ol], obtusifoliol [4 α ,14 α -dimethyl-5 α -ergosta-8,24(24¹)-dien-3 β -ol] and gramisterol [4 α -methyl-5 α -ergosta-7,24(24¹)-dien-3 β -ol] showed potential to yield information about the abundance of sedge roots and mosses (Ronkainen et al., 2013). Even though sedge root remains and mosses were present in the studied peat, the above mentioned group-specific sterols were not detected and only brassicasterol, campesterol, stigmasterol, β -sitosterol that are common to most plant species were found from fossil peat material. We attribute this absence of the plant-specific sterols is due to degradation of these compounds (Lehtonen and Ketola 1993) given that their concentrations in fen plants was rather low (Ronkainen et al., 2013), and we conclude that appearance of stanols (Fig 6) indicates degradation of organic matter since deposition (cf. McClymont et al., under review). In both cores the fen-bog transition and the layers below the transition were characterized by the presence of stanols and it seems that especially the 5 α (H)-stanols/ Δ^5 -sterol ratio, which is related to anoxic conditions and decay of plant material (Bertrand et al., 2012; McClymont et al. under review) is a strong marker for degradation, and changes in the ratio were consistent with degradation measures presented here (Fig 9). Similarly to Andersson and Meyers (2012), the CPI value increases up-core in both of the studied cores, indicating better preservation of organic matter at the top layers and a progressive degradation down core. In both cores CPI reaches its minimum right below the transition layer, almost simultaneously where the C/N ratio decreases to its minimum. In both cores the amount of UOM is clearly higher at and below the fen-bog transition than in the bog peat layers. Changes in the degradation measures might indicate that drier periods with lower water table level triggered the fen-bog transition (Hughes and Barber, 2003; Hughes, 2000), resulting in accelerated plant litter decay. This

interpretation would correspond to McClymont et al. (under review) results where high $5\alpha(\text{H})$ -stanols/ Δ^5 -sterol ratio occurred simultaneously with low water table level. Also, around the fen-bog transition layer the concentrations of sterols and triterpenoids were high in the peat, while the dominating macrofossils were sedge roots and other parts of sedges. Previous studies have stated that sterol and triterpenoid concentrations are higher in vascular plants than mosses (e.g. Pancost et al., 2002; Ronkainen et al., 2013). The results presented here support this interpretation, and suggest that these compounds originate from vascular plants. This result is potentially important because high proportions of highly decomposed organic matter hampers reliable environmental reconstructions (cf. Ficken et al., 1998), including identifying the timing of fen-bog transitions in peat cores.

Conclusions

In this study we investigated whether biomarkers can be applied to distinguish fen and bog environments, and if plant-specific biomarkers can be identified from fen peat. Not surprisingly, the palaeobotanical analyses were able to clearly separate dry bog hummocks, moist lawns and wet fen habitats apart. With the biomarkers more robust conclusions could be drawn only when the biomarkers were combined with the macrofossil data, after which a similar kind of sub-division of peatland habitats was achieved as yielded by palaeobotanical analyses. In agreement with our previous study of fen plants, we confirm that using biomarker data from highly humified fen peat layers to achieve species level information of past plant assemblages is very challenging. Although we previously showed that certain sterols could be used as indicators for some plant groups (e.g. *Sphagnum* mosses or sedge roots), these signals were not translated into the highly humified peat. However, we were able to separate bog and fen habitats apart by the changes in *n*-alkane concentrations and the *n*-alkane ratios along the cores. Moreover, the transition zone between fen and bog habitats was

characterized by high concentrations of sterols and triterpenoids originating from vascular plants. This proxy result seems to be applicable when reconstructing dominating plant groups during the highly humified peat phases, and may potentially also be used as a degradation measure as related to past changes in the water table level, and the following increase in level of decay as indicated here by the $5\alpha(\text{H})$ -stanols/ Δ^5 -sterol ratio, CPI-value, C/N ratio and high UOM.

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610 **Table caption:**

611 Table 1. Macrofossil and biomarker communities in peat cores SJ5 and SJ6 derived from
612 TWINSPAN (n=72, macrofossils and biomarkers in the data). Five cut levels and two
613 divisions were used.

614 **Supplementary data table 1. File includes published distributions of biomarkers in**
615 **plants mentioned in the study.**

616 **Supplementary data table 2. File includes all the biomarker data used in the analysis.**



Fig.1. Location of the study site. Samples were collected from two closely situated peatlands from the Siikajoki commune ((64°45'N, 24°42'E), Finland, Northern-Europe.
86x93mm (600 x 600 DPI)

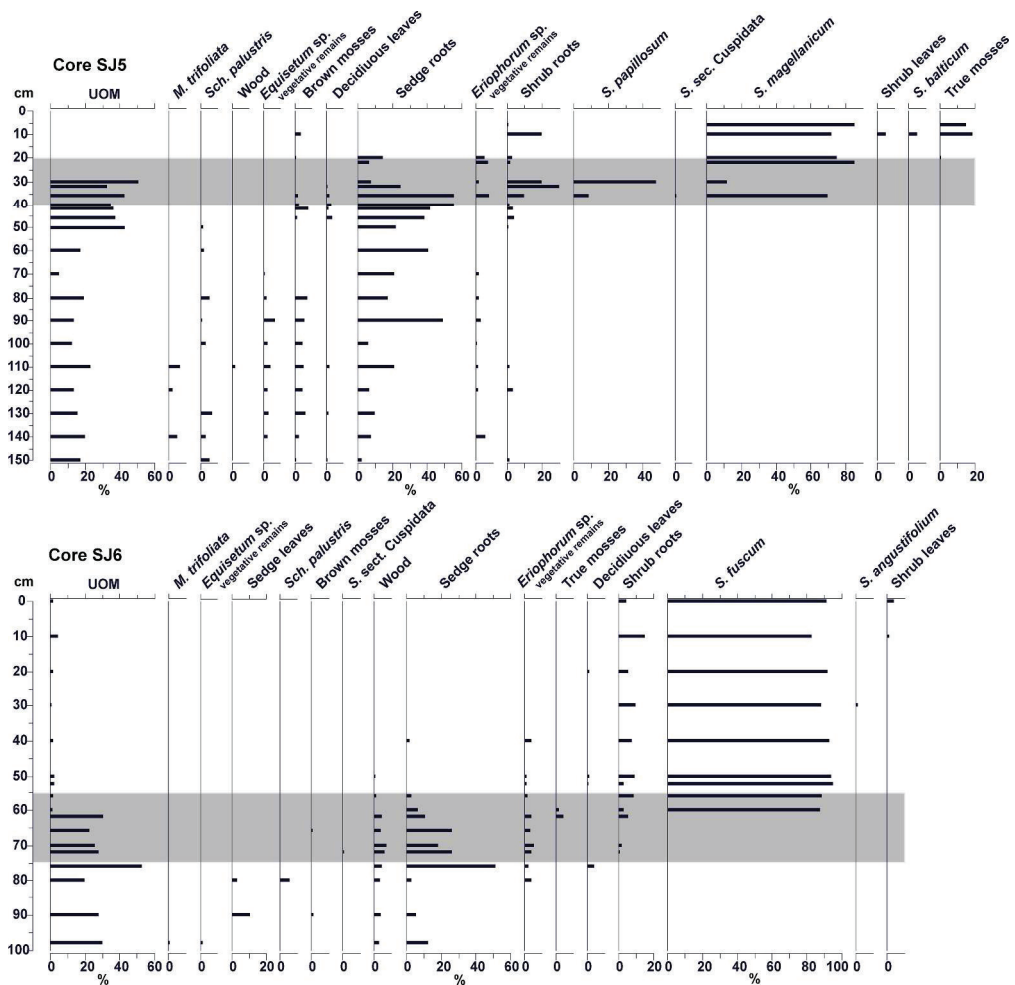


Fig. 2. Plant macrofossil records of cores SJ5 and SJ6. Macrofossil abundances are expressed as proportions (%). The fen-bog transition zone is marked with gray bar.

200x195mm (600 x 600 DPI)



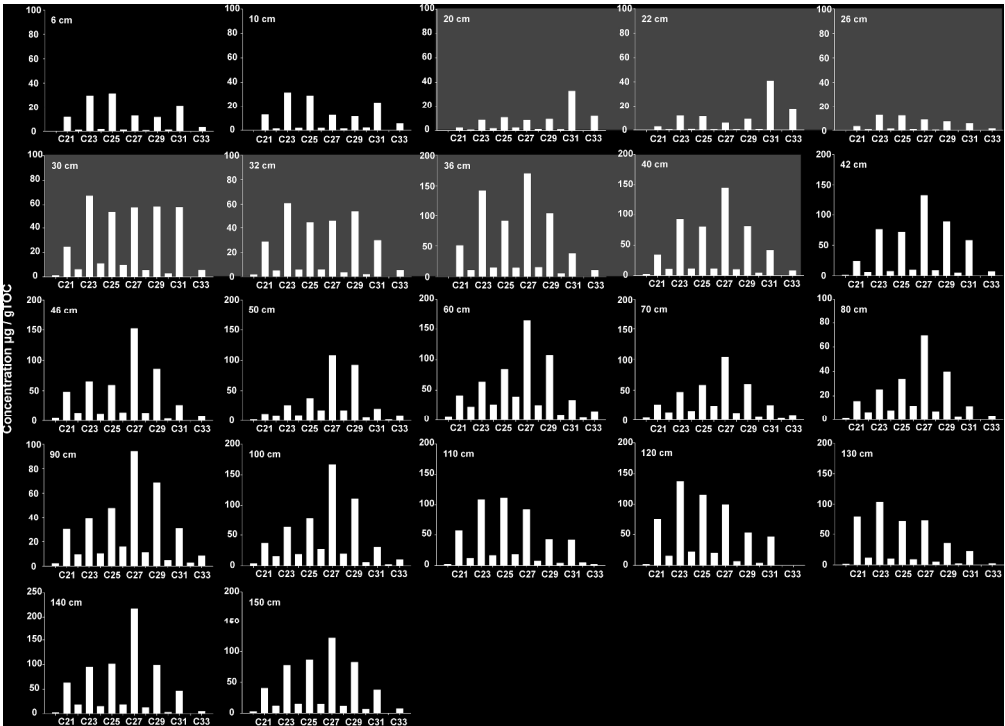


Fig. 4a. *n*-alkane concentrations ($\mu\text{g/gTOC}$) of core SJ6 by depth. The fen-bog transition zone is marked with gray.

191x137mm (600 x 600 DPI)

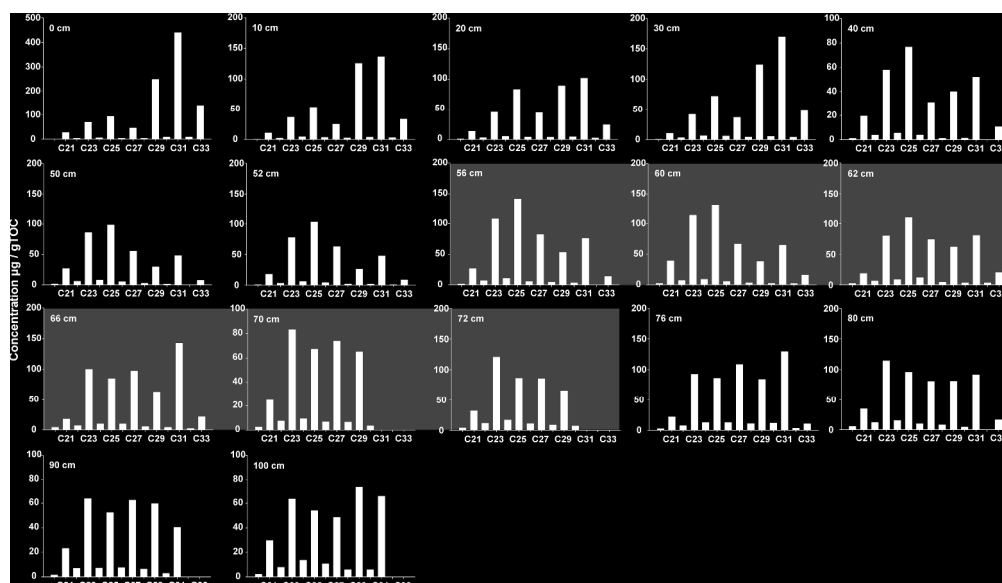


Fig. 4b. *n*-alkane concentrations ($\mu\text{g/gTOC}$) of core SJ6 by depth. The fen-bog transition zone is marked with gray.
168x96mm (600 x 600 DPI)

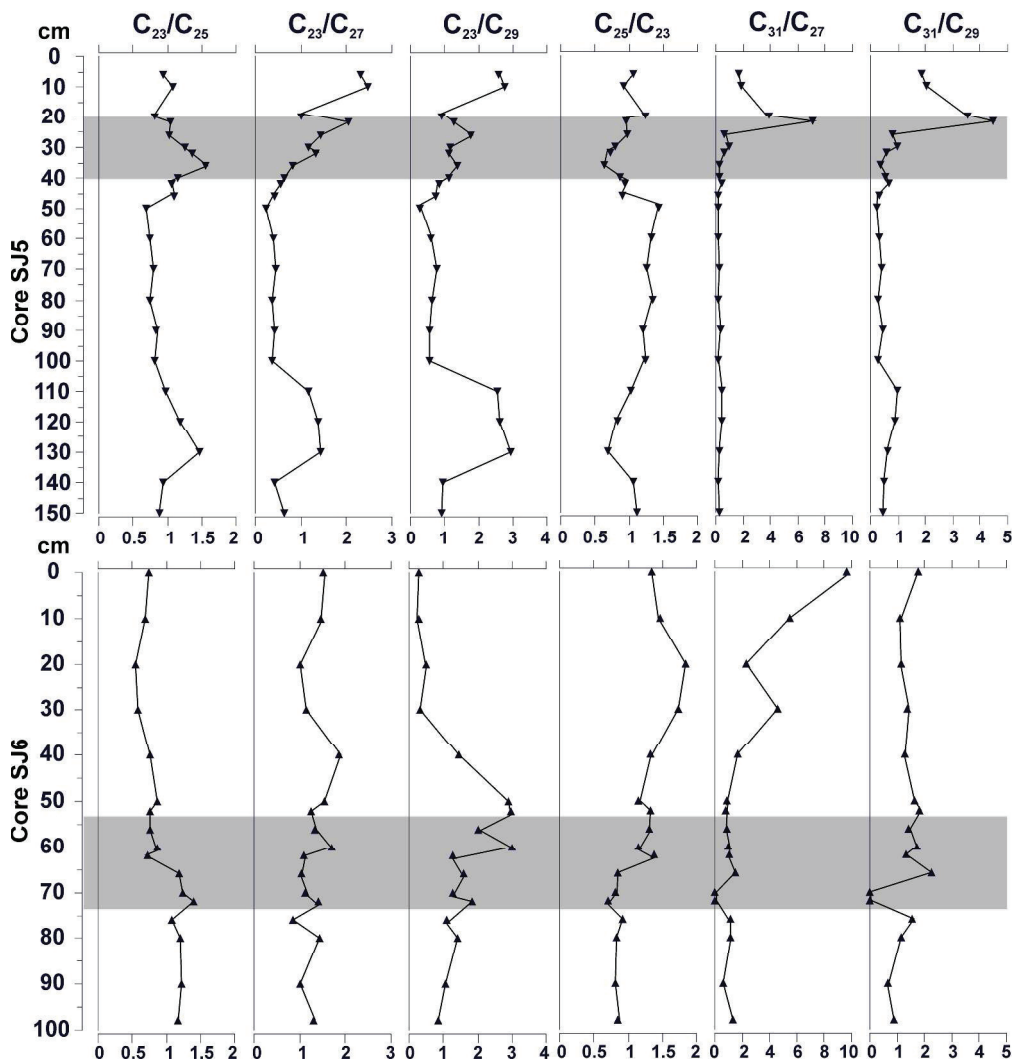


Fig. 5. *n*-alkane ratios of cores SJ5 and SJ6. The fen-bog transition zone is marked with gray bar.
188x196mm (600 x 600 DPI)

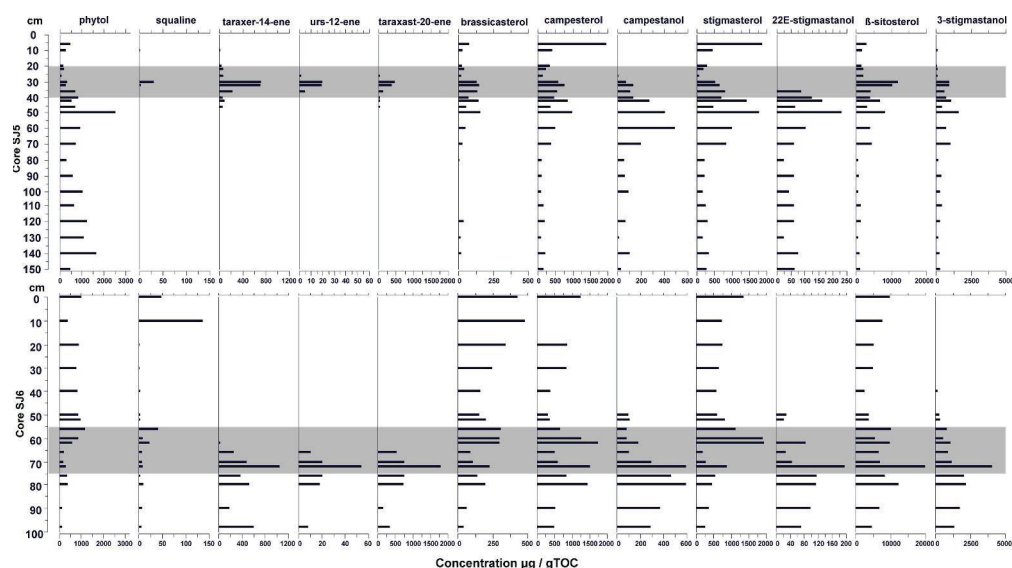


Fig. 6. Concentrations ($\mu\text{g/gTOC}$) of triterpenoids, Δ^5 -sterols and $5\alpha(\text{H})$ -stanols in cores SJ5 and SJ6. The fen-bog transition zone is marked with gray bar.
161x89mm (600 x 600 DPI)

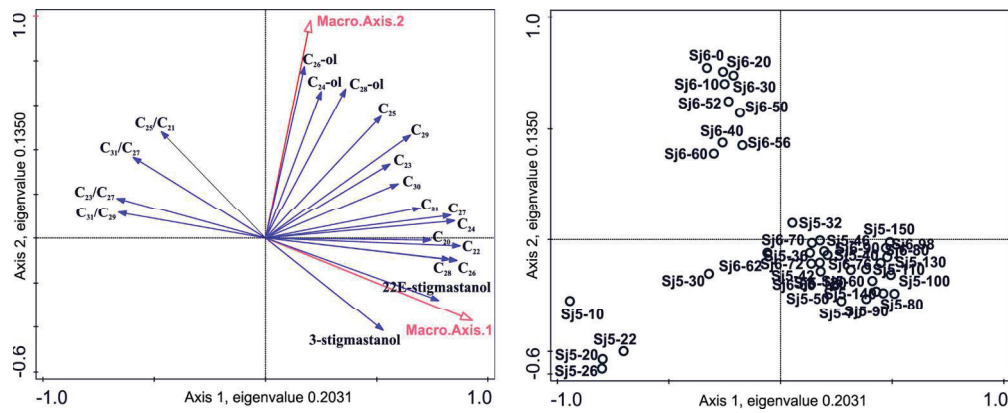


Fig. 7. RDA of biomarker data of cores SJ5 and SJ6, case scores Macro.axis.1 and Macro.axis.2 from macrofossil DCA as explanatory variables. The first axis explains 20 % (eigenvalue 0.2031) and the second axis 13 % (eigenvalue 0.1350) of the variation in the biomarker data (pseudo $F = 9.2$ and p -value = 0.002). In the figure 20 (out of 54) best fitted biomarkers are presented.

76x30mm (600 x 600 DPI)

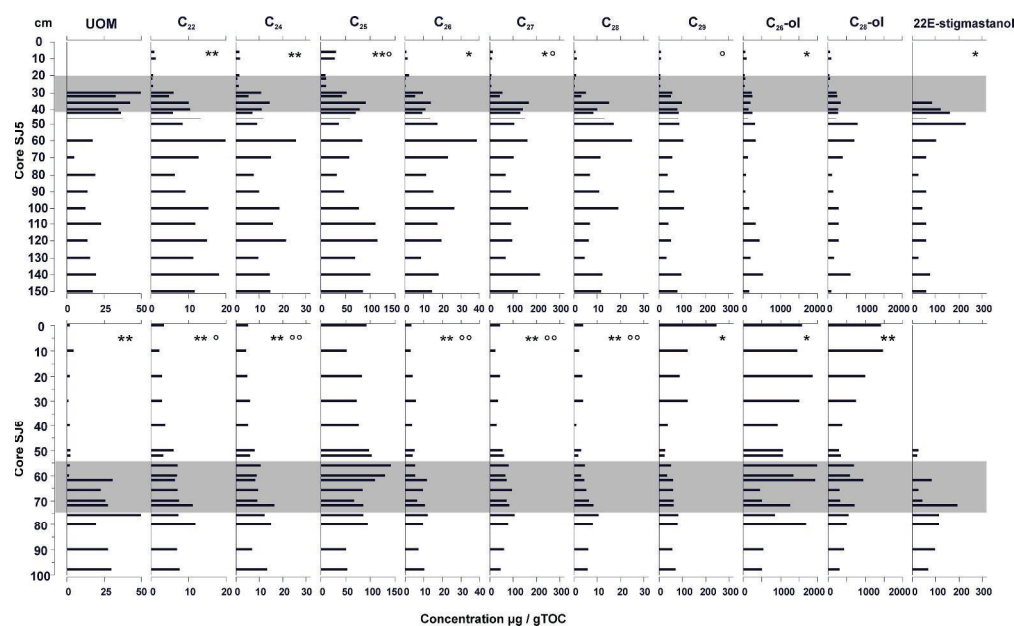


Fig. 8. Concentrations ($\mu\text{g/gTOC}$) of 10 best fitted biomarkers from biomarker RDA in cores SJ5 and SJ6. Compounds correlating with depth (sig. 0.05=*, sig. 0.01=**) and UOM (sig. 0.05=°, sig 0.01=°°) are marked. The fen-bog transition zone in both cores is marked with gray.

173x105mm (600 x 600 DPI)

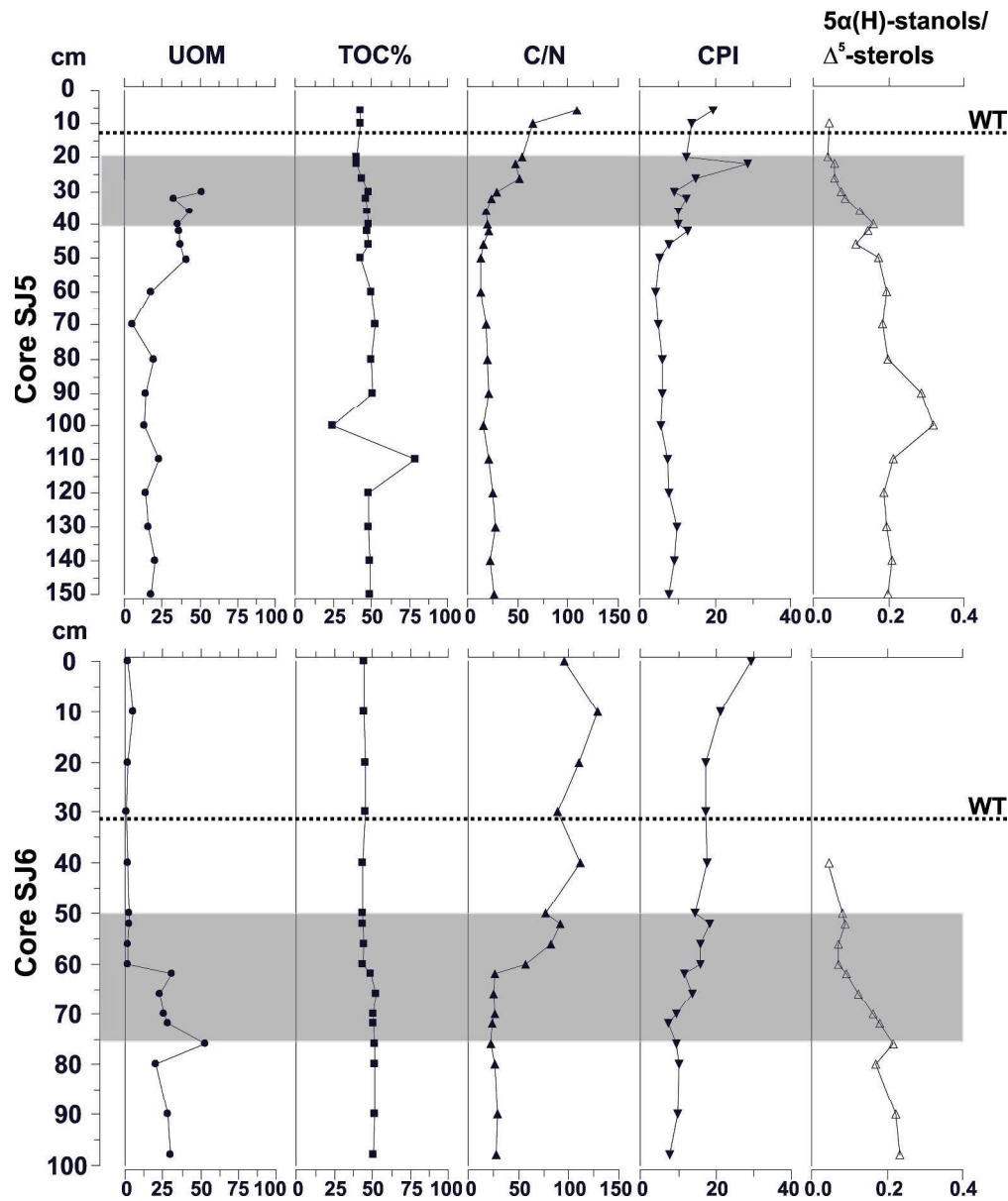


Fig. 9. Degradation measures in cores SJ5 and SJ6. UOM and TOC are presented as %, C/N ratio, CPI-value ($= \text{CPI}_{\text{alk}} = ((C_{21}-C_{31}) + (C_{23}-C_{33})\text{odd}) / 2 * (C_{22}-C_{32})\text{even}$) according to Andersson and Mayers, 2012) and the $5\alpha(\text{H})$ -stanols/ Δ^5 -sterol ratio ($= (\text{campestanol} + 22\text{E-stigmastanol} + 3\text{-stigmastanol}) / (\text{campesterol} + \text{campestanol} + \text{stigmastanol} + 22\text{E-stigmastanol} + \beta\text{-sitosterol} + 3\text{-stigmastanol})$) according to McClymony et al., 2013) The fen-bog transition zone is marked with gray bar and water table level (WT) with dashed line.
169x201mm (600 x 600 DPI)

First division	Second division
Hummock-fen $n = 33$	Hummock $n = 9$
<i>S. fuscum</i> , <i>Equisetum</i> sp.,	<i>S. fuscum</i> , shrub leaves
<i>Sch. palustris</i> , wood,	C_{27} -ol, C_{34} , squaline
deciduous leaves, UOM	
campestanol, C_{20} -ol	Fen $n = 24$
22E-stigmastanol	brown mosses, sedge roots
C_{27} , C_{20} , C_{32}	<i>Equisetum</i> sp., <i>Sch. palustris</i>
	<i>M. trifoliata</i> , campestanol,
	22E-stigmastanol, urs-20-ene
Lawn $n = 6$	
<i>S. magellanicum</i> , <i>S. papillosum</i>	
<i>S. balticum</i> , <i>S.</i> sect. <i>Cuspidata</i>	
true-mosses, urs-20-ene	

Table 1

Plant	Major homologue associated to plant	Studied plant species
<i>S. fuscum</i>	$n-C_{25}^{e, b, i, k, m}$	
<i>S. angustifolium</i>	$n-C_{23}^l$	
<i>S. magellanicum</i>	$n-C_{25}^l$	
<i>S. papillosum</i>	$n-C_{23}^l, n-C_{25}^l$	
<i>S. baliticum</i>	$n-C_{23}^{l, m}$	
<i>S. sec. Cuspidata</i>	$n-C_{23}^{m, o}$	<i>S. baliticum</i> ^m , <i>S. lindbergii</i> ^m , <i>S. angustifolium</i> ^l , <i>S. cuspidatum</i> ^l , <i>S. majus</i> ^l , <i>S. tenellum</i> ^l , <i>S. riparium</i> ^o
True mosses	$n-C_{27}^h, n-C_{31}^{e, m}$	<i>Polytrichum</i> sp. ^{e, h} , <i>Dicranum elongatum</i> ^m
Brown mosses	$n-C_{25}^o, n-C_{27}^o$	<i>Warnstorfia exannulata</i> ^o
<i>M. trifoliata</i> leaves	$n-C_{29}^o$	
<i>M. trifoliata</i> roots	$n-C_{21}^o, n-C_{23}^o$	
<i>Equisetum</i> sp.	no reference	
<i>S. palustris</i>	no reference	
<i>Eriophorum</i> sp. leaves	$n-C_{27}^{o, q}, n-C_{29}^m$	
<i>Eriophorum</i> sp roots	$n-C_{27}^{o, m, o}$	
Sedge leaves	$n-C_{27}^{b, m}, n-C_{29}^{f, o}$ or $n-C_{31}^{f, o}$	<i>Carex</i> sp. ^{b, f, m, o}
Sedge roots	$n-C_{21}^o, n-C_{23}^{m, o}, n-C_{27}^o$	
Shrub leafs	$n-C_{27}^{d, i, q}, n-C_{29}^{b, q}, n-C_{31}^{a, q}$	<i>Ledum</i> sp. ^{a, m} , <i>V. vitis-idaea</i> ^b , <i>B. nana</i> ^{d, j, m} , <i>E. nigrum</i> ^m , <i>V. uliginosum</i> ^m , several species ^q
Shrub root	$n-C_{27}^m, n-C_{29}^m, n-C_{31}^n$	<i>Eriaceaceae</i> ⁿ , <i>Betula nana</i> ^{n, m} , <i>L. palustris</i> ^m , <i>E. nigrum</i> ^m , <i>V. uliginosum</i> ^m
Wood	$n-C_{27}^m$	<i>Betula</i> (tree) ^m
Deciduous leaves	$n-C_{27}^{m, q}, n-C_{29}^{m, q}$	<i>Betula</i> (tree) ^m , several species ^q
Conifer needles	$n-C_{27}^q, n-C_{29}^q, n-C_{31}^q$	several species ^q

^a (Salasoo, 1987); ^b (Ficken et al., 1998); ^c (Corrigan et al., 1973); ^d (Sachse et al., 2006); ^e (Nott et al., 2000); ^f (Ficken et al., 2000); ^g (Baas et al., 2000); ^h (Nissinen and Sewón, 1994); ⁱ (Vonk and Gustafsson, 2009); ^j (Zech et al., 2010); ^k (Dembitsky, 1993); ^l (Bingham et al., 2010); ^m (Ronkainen et al. submitted); ⁿ (Andersson et al., 2011); ^o (Ronkainen et al. 2013); ^q (Tarasov et al., 2013)

n-alkanes

Sample	C20	C21	C22	C23	C24	C25	C26	C27	C28	C29	C30	C31	C32	C33	C34	C35
SI5: 6	-	11.5	1.1	29.6	1.7	31.5	1.2	12.8	0.8	11.4	1.2	21.4	-	3.3	-	-
SI5: 10	-	12.9	1.4	31.5	1.9	29.0	1.9	12.7	1.5	11.4	2.2	23.1	-	5.5	-	-
SI5: 20	-	2.3	0.7	8.7	1.7	10.7	2.3	8.6	1.0	9.3	0.8	33.0	-	11.9	-	-
SI5: 22	-	3.0	0.5	11.8	0.9	11.3	0.5	5.8	0.5	9.2	0.7	41.0	-	17.1	-	-
SI5: 26	-	3.2	0.8	12.6	1.3	12.2	0.8	8.8	0.6	7.1	-	5.8	-	1.5	-	-
SI5: 30	1.2	24.7	6.2	67.2	11.0	53.2	9.8	57.6	5.4	58.2	2.8	57.8	-	5.7	-	-
SI5: 32	1.7	28.6	4.9	60.8	5.6	44.3	5.7	45.5	3.4	53.2	1.8	29.6	-	5.2	-	-
SI5: 36	-	50.8	10.3	142.0	14.7	91.0	14.2	170.2	15.2	103.3	5.0	37.7	-	10.3	-	-
SI5: 40	2.3	33.9	10.6	92.2	11.3	80.0	11.3	145.2	10.1	80.8	4.5	41.4	-	8.1	-	0.7
SI5: 42	1.2	24.1	6.0	76.0	7.4	71.7	9.5	133.1	8.6	88.7	4.9	58.0	-	7.1	-	-
SI5: 46	4.2	48.0	13.3	64.9	12.0	59.1	13.8	152.1	13.3	85.7	3.3	26.4	-	6.7	-	-
SI5: 50	1.5	11.6	8.6	26.1	9.3	37.4	17.6	108.0	17.3	92.0	4.8	20.0	1.5	9.1	-	1.4
SI5: 60	4.9	41.0	22.3	63.5	26.1	84.4	39.0	164.1	25.2	107.1	9.2	33.4	3.9	14.9	-	1.3
SI5: 70	3.2	25.7	12.9	46.2	15.2	57.9	23.4	103.9	11.6	59.1	4.5	24.8	2.7	8.4	-	-
SI5: 80	1.2	15.6	6.5	25.1	7.9	33.8	11.7	69.5	7.1	39.8	2.1	11.3	-	3.2	-	-
SI5: 90	2.2	30.0	9.3	39.4	10.2	47.7	15.7	94.3	11.0	68.6	4.6	30.5	2.8	8.3	1.0	1.6
SI5: 100	3.7	37.0	15.5	63.2	19.0	78.4	27.1	166.7	19.4	110.8	5.6	30.5	2.3	10.3	-	1.9
SI5: 110	2.1	56.6	11.9	108.6	16.3	111.4	17.9	92.5	7.1	42.4	3.9	41.7	4.8	1.9	-	-
SI5: 120	1.6	74.7	15.0	137.2	21.9	115.1	20.0	99.5	6.5	52.7	3.6	46.2	-	-	-	-
SI5: 130	1.5	79.8	11.4	103.5	10.0	70.6	8.9	72.0	4.9	35.3	1.9	22.1	-	2.4	-	-
SI5: 140	2.0	62.5	18.2	94.6	14.8	101.0	18.5	217.3	12.4	98.6	2.9	46.2	-	4.4	-	-
SI5: 150	2.6	40.6	11.8	77.6	15.0	86.9	14.7	122.1	11.8	82.8	6.6	37.7	-	7.4	-	-

Sample	C20	C21	C22	C23	C24	C25	C26	C27	C28	C29	C30	C31	C32	C33	C34	C35
SI6: 0	0.6	27.8	3.5	69.4	5.3	93.4	3.7	45.8	4.1	248.6	8.5	439.9	8.8	140.2	0.8	-
SI6: 10	0.3	10.8	2.3	36.4	4.5	53.3	3.3	24.9	2.4	125.5	3.6	136.2	3.0	33.1	-	-
SI6: 20	1.0	13.5	2.9	45.0	5.2	83.0	4.1	44.1	3.6	89.3	4.4	101.8	2.4	24.4	0.6	-
SI6: 30	0.5	10.3	2.9	41.7	6.2	72.4	5.8	36.5	4.0	124.0	5.3	169.7	3.9	48.0	-	-
SI6: 40	1.1	19.5	3.9	57.9	5.5	76.7	3.9	31.1	1.1	40.1	1.4	51.9	-	10.6	-	-
SI6: 50	1.8	26.8	6.1	85.7	8.1	98.2	5.3	55.4	3.0	29.6	1.6	48.2	-	7.4	-	-
SI6: 52	0.7	17.8	3.3	77.6	6.2	103.0	4.2	62.7	2.1	26.1	1.6	47.6	0.9	8.6	-	-
SI6: 56	1.8	26.8	7.1	108.0	10.7	141.5	5.7	82.1	4.7	53.2	3.2	75.7	-	13.7	-	-
SI6: 60	2.4	39.1	7.0	113.9	9.0	130.4	5.5	66.7	3.0	38.0	2.4	64.8	1.9	15.9	-	-
SI6: 62	2.4	18.8	6.4	79.8	8.5	109.8	11.9	74.0	4.5	61.9	3.2	80.8	3.3	20.1	-	-
SI6: 66	4.2	17.8	7.2	100.1	9.6	84.9	9.7	97.2	5.4	63.1	4.1	142.5	2.1	21.3	-	-
SI6: 70	2.5	25.1	7.6	83.2	9.4	67.2	6.8	73.9	6.6	65.0	3.5	-	-	-	-	-
SI6: 72	3.6	31.9	11.2	120.5	16.6	85.8	10.9	85.3	8.6	65.5	6.9	-	-	-	-	-
SI6: 76	2.4	21.6	7.4	92.5	12.4	85.8	12.4	108.5	10.7	83.9	11.3	129.4	3.2	10.4	-	-
SI6: 80	5.2	34.6	11.9	114.3	15.2	95.4	9.8	80.3	8.1	80.8	4.4	91.4	-	16.5	-	-
SI6: 90	1.5	23.2	7.0	63.7	7.2	52.4	7.5	62.4	6.3	59.7	2.9	40.3	-	-	-	-
SI6: 98	2.1	29.6	7.8	63.5	13.7	54.0	10.6	48.5	5.8	73.9	5.8	65.7	-	-	-	-

Concentrations are presented as µg gTOC

n-alkane ratios

Sample	C23/C25	C23/C27	C23/C29	C23/C31	C25/C29	C31/C27	C31/C29	C33/C31	C23/(C23+C29)
SI5: 6	0.9	2.3	2.6	1.4	2.8	1.7	1.9	0.2	0.7
SI5: 10	1.1	2.5	2.8	1.4	2.5	1.8	2.0	0.2	0.7
SI5: 20	0.8	1.0	0.9	0.3	1.2	3.9	3.6	0.4	0.5
SI5: 22	1.0	2.0	1.3	0.3	1.2	7.0	4.5	0.4	0.6
SI5: 26	1.0	1.4	1.8	2.2	1.7	0.7	0.8	0.3	0.6
SI5: 30	1.3	1.2	1.2	1.2	0.9	1.0	1.0	0.1	0.5
SI5: 32	1.4	1.3	1.1	2.1	0.8	0.7	0.6	0.2	0.5
SI5: 36	1.6	0.8	1.4	3.8	0.9	0.2	0.4	0.3	0.6
SI5: 40	1.2	0.6	1.1	2.2	1.0	0.3	0.5	0.2	0.5
SI5: 42	1.1	0.6	0.9	1.3	0.8	0.4	0.7	0.1	0.5
SI5: 46	1.1	0.4	0.8	2.5	0.7	0.2	0.3	0.3	0.4
SI5: 50	0.7	0.2	0.3	1.3	0.4	0.2	0.2	0.5	0.2
SI5: 60	0.8	0.4	0.6	1.9	0.8	0.2	0.3	0.4	0.4
SI5: 70	0.8	0.4	0.8	1.9	1.0	0.2	0.4	0.3	0.4
SI5: 80	0.7	0.4	0.6	2.2	0.8	0.2	0.3	0.3	0.4
SI5: 90	0.8	0.4	0.6	1.3	0.7	0.3	0.4	0.3	0.4
SI5: 100	0.8	0.4	0.6	2.1	0.7	0.2	0.3	0.3	0.4
SI5: 110	1.0	1.2	2.6	2.6	2.6	0.5	1.0	0.0	0.7
SI5: 120	1.2	1.4	2.6	3.0	2.2	0.5	0.9	0.0	0.7
SI5: 130	1.5	1.4	2.9	4.7	2.0	0.3	0.6	0.1	0.7
SI5: 140	0.9	0.4	1.0	2.0	1.0	0.2	0.5	0.1	0.5
SI5: 150	0.9	0.6	0.9	2.1	1.0	0.3	0.5	0.2	0.5

Sample	C23/C25	C23/C27	C23/C29	C23/C31	C25/C29	C31/C27	C31/C29	C33/C31	C23/(C23+C29)
SI6: 0	0.7	1.5	0.3	0.2	0.4	9.6	1.8	0.3	0.2
SI6: 10	0.7	1.5	0.3	0.3	0.4	5.5	1.1	0.2	0.2
SI6: 20	0.5	1.0	0.5	0.4	0.9	2.3	1.1	0.2	0.3
SI6: 30	0.6	1.1	0.3	0.2	0.6	4.6	1.4	0.3	0.3
SI6: 40	0.8	1.9	1.4	1.1	1.9	1.7	1.3	0.2	0.6
SI6: 50	0.9	1.5	2.9	1.8	3.3	0.9	1.6	0.2	0.7
SI6: 52	0.8	1.2	3.0	1.6	3.9	0.8	1.8	0.2	0.7
SI6: 56	0.8	1.3	2.0	1.4	2.7	0.9	1.4	0.2	0.7
SI6: 60	0.9	1.7	3.0	1.8	3.4	1.0	1.7	0.2	0.7
SI6: 62	0.7	1.1	1.3	1.0	1.8	1.1	1.3	0.2	0.6
SI6: 66	1.2	1.0	1.6	0.7	1.3	1.5	2.3	0.1	0.6
SI6: 70	1.2	1.1	1.3	-	1.0	0.0	0.0	-	0.6
SI6: 72	1.4	1.4	1.8	-	1.3	0.0	0.0	-	0.6
SI6: 76	1.1	0.9	1.1	0.7	1.0	1.2	1.5	0.1	0.5
SI6: 80	1.2	1.4	1.4	1.3	1.2	1.1	1.1	0.2	0.6
SI6: 90	1.2	1.0	1.1	1.6	0.9	0.6	0.7	0.0	0.5
SI6: 98	1.2	1.3	0.9	1.0	0.7	1.4	0.9	0.0	0.5

n-alkane ratios

Sample	C25/(C25+C29)	C23/(C27+C31)	Paq	ACL C17-C35	Pwax	C23/C21	C21/C23	C25/C21	C25/C23
SI5: 6	1.0	0.9	0.7	26.0	0.4	2.6	0.4	2.7	1.1
SI5: 10	0.9	0.9	0.6	26.1	0.4	2.4	0.4	2.3	0.9
SI5: 20	0.8	0.2	0.3	28.8	0.7	3.8	0.3	4.6	1.2
SI5: 22	1.0	0.3	0.3	29.0	0.7	4.0	0.3	3.8	1.0
SI5: 26	0.9	0.9	0.7	26.1	0.5	3.9	0.3	3.8	1.0
SI5: 30	0.8	0.6	0.5	26.5	0.6	2.7	0.4	2.2	0.8
SI5: 32	0.9	0.8	0.6	26.0	0.5	2.1	0.5	1.5	0.7
SI5: 36	0.9	0.7	0.6	25.9	0.6	2.8	0.4	1.8	0.6
SI5: 40	0.9	0.5	0.6	26.2	0.6	2.7	0.4	2.4	0.9
SI5: 42	0.9	0.4	0.5	26.7	0.7	3.2	0.3	3.0	0.9
SI5: 46	0.8	0.4	0.5	26.1	0.7	1.4	0.7	1.2	0.9
SI5: 50	0.7	0.2	0.4	27.2	0.8	2.3	0.4	3.2	1.4
SI5: 60	0.7	0.3	0.5	26.5	0.7	1.5	0.6	2.1	1.3
SI5: 70	0.7	0.4	0.6	26.3	0.6	1.8	0.6	2.2	1.3
SI5: 80	0.7	0.3	0.5	26.3	0.7	1.6	0.6	2.2	1.3
SI5: 90	0.8	0.3	0.5	26.5	0.7	1.3	0.8	1.6	1.2
SI5: 100	0.7	0.3	0.5	26.5	0.7	1.7	0.6	2.1	1.2
SI5: 110	0.9	0.8	0.7	25.4	0.4	1.9	0.5	2.0	1.0
SI5: 120	0.9	0.9	0.7	25.2	0.4	1.8	0.5	1.5	0.8
SI5: 130	0.9	1.1	0.8	24.7	0.4	1.3	0.8	0.9	0.7
SI5: 140	0.8	0.4	0.6	26.1	0.6	1.5	0.7	1.6	1.1
SI5: 150	0.9	0.5	0.6	26.1	0.6	1.9	0.5	2.1	1.1

Sample	C25/(C25+C29)	C23/(C27+C31)	Paq	ACL C17-C35	Pwax	C23/C21	C21/C23	C25/C21	C25/C23
SI6: 0	0.3	0.1	0.2	29.3	1.1	2.5	0.4	3.4	1.3
SI6: 10	0.3	0.2	0.2	28.6	1.1	3.4	0.3	4.9	1.5
SI6: 20	0.5	0.3	0.4	27.8	0.9	3.3	0.3	6.1	1.8
SI6: 30	0.4	0.2	0.3	28.7	1.0	4.1	0.2	7.1	1.7
SI6: 40	0.7	0.7	0.5	26.4	0.6	3.0	0.3	3.9	1.3
SI6: 50	0.8	0.8	0.6	25.8	0.5	3.2	0.3	3.7	1.1
SI6: 52	0.8	0.7	0.6	26.0	0.5	4.4	0.2	5.8	1.3
SI6: 56	0.7	0.7	0.5	26.2	0.5	4.0	0.2	5.3	1.3
SI6: 60	0.8	0.9	0.6	25.8	0.5	2.9	0.3	3.3	1.1
SI6: 62	0.6	0.5	0.5	26.7	0.6	4.3	0.2	5.8	1.4
SI6: 66	0.6	0.4	0.4	27.2	0.7	5.6	0.2	4.8	0.8
SI6: 70	0.5	1.1	0.5	25.3	0.6	3.3	0.3	2.7	0.8
SI6: 72	0.6	1.4	0.6	25.1	0.5	3.8	0.3	2.7	0.7
SI6: 76	0.5	0.4	0.4	27.1	0.8	4.3	0.2	4.0	0.9
SI6: 80	0.5	0.7	0.5	26.4	0.7	3.3	0.3	2.8	0.8
SI6: 90	0.5	0.6	0.4	26.1	0.7	2.7	0.4	2.3	0.8
SI6: 98	0.4	0.6	0.4	26.4	0.8	2.1	0.5	1.8	0.8

Triterpenoids

Sample	squalene	taraxer-14-ene	urs-12-ene	taraxast-20-ene
SI5: 6	-	1	-	-
SI5: 10	2.7	18.6	-	3.3
SI5: 20	-	37.4	-	2.1
SI5: 22	-	78.4	-	6.7
SI5: 26	-	69.0	2.118735599	57.6
SI5: 30	31.6	723.1	20.2	475.0
SI5: 32	3.4	713.9	19.8	386.8
SI5: 36	-	229.4	5.4	145.8
SI5: 40	-	59.2	-	50.6
SI5: 42	-	100.4	-	47.0
SI5: 46	-	66.9	-	46.2
SI5: 50	-	8.9	-	-
SI5: 60	-	5.3	-	-
SI5: 70	-	3.6	-	1.1
SI5: 80	-	3.8	-	-
SI5: 90	-	4.4	-	0.9
SI5: 100	-	7.3	-	3.3
SI5: 110	-	-	-	-
SI5: 120	-	-	-	-
SI5: 130	-	-	-	-
SI5: 140	-	-	-	-
SI5: 150	-	-	-	-

Sample	squalene	taraxer-14-ene	urs-12-ene	taraxast-20-ene
SI6: 0	48.2	-	-	1.8
SI6: 10	137.3	3.1	-	-
SI6: 20	2.5	4.1	-	-
SI6: 30	2.4	-	-	-
SI6: 40	3.3	-	-	-
SI6: 50	3.6	10.3	-	9.6
SI6: 52	3.6	8.4	-	4.7
SI6: 56	41.5	-	-	-
SI6: 60	9.3	2.9	-	2.6
SI6: 62	23.1	27.1	-	24.8
SI6: 66	7.9	263.4	10.4	546.5
SI6: 70	8.0	484.1	20.8	767.6
SI6: 72	8.7	1053.5	54.1	1805.1
SI6: 76	3.8	383.7	20.6	755.0
SI6: 80	10.0	526.6	18.7	745.1
SI6: 90	7.9	190.5	-	166.3
SI6: 98	6.6	602.3	20.3	57.9

Concentrations are presented as $\mu\text{g g}^{-1}\text{TOC}$

Sterols					Stanols				
Sample	phytol	α -tocopherol	brassicasterol	campesterol	stigmasterol	β -sitosterol	campestanol	22E-stigmasterol-22-en-3 β -ol	3-stigmastanol
SJ5: 6	488.1	-	80.6	1959.4	1873.2	3065.0	-	-	-
SJ5: 10	287.9	-	31.5	427.1	465.2	1770.2	-	-	115.5
SJ5: 20	176.9	-	26.1	362.3	304.3	1609.8	-	-	92.4
SJ5: 22	185.6	-	42.3	243.7	189.9	2050.1	-	-	146.4
SJ5: 26	94.8	-	21.4	150.9	75.2	2035.8	11.8	-	126.2
SJ5: 30	336.3	-	134.9	596.3	539.7	12073.3	73.4	-	991.1
SJ5: 32	274.3	-	149.0	773.1	652.0	10419.6	136.6	-	959.5
SJ5: 36	718.9	-	138.1	562.1	820.9	4251.1	111.6	88.6	605.4
SJ5: 40	858.0	-	77.0	473.7	700.4	4134.5	136.3	125.3	743.1
SJ5: 42	537.6	-	146.2	874.8	1438.0	6869.6	275.3	164.1	1105.6
SJ5: 46	713.8	-	57.1	374.2	472.9	3245.9	-	65.3	462.7
SJ5: 50	2537.7	-	158.8	984.2	1779.8	8246.0	408.1	233.1	1651.2
SJ5: 60	946.6	-	54.9	507.6	1003.9	3984.2	491.2	103.4	746.4
SJ5: 70	730.6	-	30.3	382.9	840.9	4538.4	203.5	61.7	1046.4
SJ5: 80	322.6	-	7.2	124.1	229.8	751.9	56.7	27.6	187.2
SJ5: 90	606.1	-	-	130.7	226.3	948.7	66.2	61.5	394.1
SJ5: 100	1039.8	-	4.6	112.2	182.2	705.8	96.6	45.0	329.8
SJ5: 110	651.9	-	-	178.0	263.0	1407.7	-	62.7	431.8
SJ5: 120	1241.6	-	38.3	214.2	314.1	1445.4	70.1	62.7	324.7
SJ5: 130	1094.2	-	17.9	107.2	173.3	620.1	14.9	27.3	175.5
SJ5: 140	1675.5	-	23.9	222.7	352.9	1134.8	104.4	77.9	271.0
SJ5: 150	485.5	-	-	179.5	290.1	1153.2	31.1	63.5	307.1

Sample	phytol	α -tocopherol	brassicasterol	campesterol	stigmasterol	β -sitosterol	campestanol	22E-stigmasterol-22-en-3 β -ol	3-stigmastanol
SJ6: 0	1031.1	-	428.7	1257.5	1363.7	9863.4	-	-	-
SJ6: 10	397.0	-	481.0	-	738.9	7705.1	-	-	-
SJ6: 20	906.8	-	343.6	862.7	753.4	5262.0	-	-	-
SJ6: 30	783.1	-	246.6	840.9	660.4	5171.0	-	-	-
SJ6: 40	839.0	-	164.4	394.8	580.2	2604.4	-	-	165.8
SJ6: 50	864.6	-	153.6	318.1	610.1	3965.4	99.0	29.5	295.2
SJ6: 52	997.1	-	204.4	375.8	838.5	3930.9	109.3	23.6	368.8
SJ6: 56	1183.2	-	309.6	675.8	1132.2	10309.6	86.3	-	851.2
SJ6: 60	881.5	-	299.2	1282.8	1911.0	5639.8	86.2	-	584.3
SJ6: 62	591.6	-	299.2	1744.0	1943.0	9975.1	186.9	85.5	1108.8
SJ6: 66	213.8	-	92.6	515.4	201.8	6789.6	107.8	28.9	914.4
SJ6: 70	208.0	-	108.9	605.6	277.9	7136.2	295.7	46.4	1198.2
SJ6: 72	299.3	-	232.2	1527.3	891.0	19970.0	596.3	196.3	4082.6
SJ6: 76	360.8	-	140.7	842.6	554.8	8407.8	466.7	116.7	2100.5
SJ6: 80	146.2	146.2	200.2	1448.1	457.6	12314.4	593.6	115.5	2204.7
SJ6: 90	146.2	205.9	67.7	529.2	375.9	6840.8	371.4	99.2	1759.6
SJ6: 98	141.9	55.7	45.0	497.6	267.0	4859.4	291.9	72.0	1360.2

Concentrations are presented as μ g gTOC